Pacific Ocean Variability in an ACCESS 1/4°-ocean Coupled Model

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1 Past Work: Tropical Instability Waves and Mixing

2 Future Work: Pacific Ocean Variability

3 1/4°-ocean ACCESS Coupled Model: Some Initial Analysis

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ROMS Modelling of Tropical Instability Waves (TIWs)



Sea Surface Temperature (°C), year day = 259

Figure: TIWs propagating along the cold tongue in a $1/4^\circ$ ROMS model of the tropical Pacific

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Figure: TIWs propagating along the cold tongue in a $1/4^\circ$ ROMS model of the tropical Pacific

TIWs play an important role in the SST budget (both lateral and vertical fluxes). Graham (2014) [1]: TIW eddy heating 37% larger in a $1/4^{\circ}$ vs. 1° NEMO simulation. ENSO state dependent: 74% larger during La Nina/Neutral conditions.

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What role does the ocean play in decadal variability in the Pacific?

- North Pacific Rossby waves
- Gyre and WBC anomalous advection, non-eddying and eddying gyre modes
- Meridional overturning: Subtropical cell variability



Proposed experiments with both 1° and $1/4^\circ\text{-}ocean$ ACCESS CMs:

- Partial coupling: SST pacemaker in North Pacific, how does this constrain tropical variability?
- Partial blocking:

 Remove oceanic tunnel between extratropics and tropics.
 Remove the mid latitude

2) Remove the mid-latitude Rossby wave pathways?

The role of eddies in tropical Pacific overturning

What role does the eddy-driven circulation play in the tropics?



Figure: (left) Mean and (right) residual overturning streamfunctions in the tropical Pacific. From Hazeleger et. al. (2001) [2]

- TIWs are well-resolved at 1/4°, but not at 1°.
- TIWs are assymmetric w.r.t. hemisphere and ENSO events.

Need for accumulation of sub-monthly eddy fluxes.

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ACCESS 1/4° Biases: SST













Figure: Comparison of ACCESS CM and HadSST2 mean SST

Figure: Focus on key regions

22.75

16.25

ACCESS 1/4° Biases: Wind Stress



Figure: Comparison of ACCESS CM and ERA Interim wind stresses



Figure: Focus on key regions



Figure: 1-6yr RMS Variability





Figure: 1-6yr RMS Variability



Figure: 1-6yr RMS Variability



Figure: (top) Nino 3 Eastern Pacific index and (bottom) power spectra



Figure: (left) Nino 3 (East Pacific) vs. Nino 4 (Central Pacific) DJF anomalies for identified La Nina and El Nino events. (right) Similar figure from CMIP5 simulations (Taschetto et. al. (2014)) [4].



Figure: Regressions of (top) SST, (middle) wind stress and (bottom) 20° C isotherm depth anomalies onto onto Nino 3.4

Pacific Ocean Variability



Figure: (left) SST and (right) τ_x anomaly composites for El Nino.



Figure: (left) SST and (right) τ_x anomaly composites for La Nina.

ACCESS 1/4° Decadal Variability



Figure: 7-40yr band-passed RMS Variability

ACCESS 1/4° Decadal Variability



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ACCESS 1/4° Decadal Variability: PDO







Figure: ERSST 1900+ PDO

ACCESS 1/4° Decadal Variability: IPO



Figure: ACCESS 250-year IPO



Figure: ERSST 1854+ IPO

IPO tripole index of Henley et. al. (2015) [3].



Figure: ACCESS 250-year correlation between SST and IPO tripole index.

Figure: Correlation between IPO tripole index and HadISST2 (from Henley et. al. (2015)).

ACCESS 1/4° Decadal Variability: Meridional Overturning



Figure: 250-year eulerian mean meridional overturning streamfunction in the Indo-Pacific.



Figure: (top) PDO time series. (middle) Northern subtropical cell transport. (bottom) 6-year rolling-mean normalized time series.

- Mean SST biases:
 - Western boundary currents $(3 4^{\circ}C \text{ too warm})$
 - Eastern upwelling regions $(3 5^{\circ}C \text{ too warm})$
 - Southern ocean $(2 3^{\circ}C \text{ too warm})$
 - Western Equatorial Pacific $(1 2^{\circ}C \text{ too cold})$

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- ENSO variability is:
 - Too strong in the eastern Pacific
 - Too fast (2.5-4 years vs. 3-7 years)
 - Too symmetric: El Nino vs. La Nina
 - Extreme eastern Pacific La Nina's (DJF Nino 3 reaching -2.6°C). Outside observational and CMIP5 range.

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- Pacific decadal variability is:
 - Too strong in WBC region.
 - Biased to Northern Hemisphere.

- T. Graham. The importance of eddy permitting model resolution for simulation of the heat budget of tropical instability waves. *Ocean Model.*, 79(0):21 – 32, 2014.
- [2] W. Hazeleger, P. de Vries, and G. J. van Oldenborgh. Do tropical cells ventilate the indo-pacific equatorial thermocline? *Geophys. Res. Lett.*, 28(9):1763–1766, 2001.
- [3] B. J. Henley, J. Gergis, D. J. Karoly, S. Power, J. Kennedy, and C. K. Folland. A tripole index for the interdecadal pacific oscillation. *Climate Dynamics*, 45(11-12):3077–3090, 2015.
- [4] A. S. Taschetto, A. S. Gupta, N. C. Jourdain, A. Santoso, C. C. Ummenhofer, and M. H. England. Cold tongue and warm pool enso events in cmip5: Mean state and future projections. *Journal of Climate*, 27(8):2861–2885, 2014.

ACCESS 1/4° Interannual Variability: Nino 3.4 Seasonality



Figure: Nino 3.4 seasonality.

ACCESS 1/4° Interannual Variability: Mean seasonality



Figure: Seasonality of the mean SST.

ACCESS 1/4° Decadal Variability: PDO



Figure: ACCESS 250-year PDO



Figure: HadISST 1900+ PDO

ACCESS 1/4° Decadal Variability: PDO





Figure: ACCESS 250-year PDO

Figure: ORA-S4 1958+ PDO



Figure: (top) ACCESS PDO Spectra (bottom) Northern Subtropical cell transport Spectra.